Thermodynamics and log-contrast analysis in fluid geochemistry

Antonella Buccianti, Barbara Nisi and Orlando Vaselli

Dipartimento di Scienze della Terra, Università degli Studi di Firenze Firenze, Italy.

Mailing address:

A. Buccianti Dip. Scienze della Terra, Università degli Studi di Firenze Via G. La Pira 4, 50125, Firenze, Italy e-mail:antonella.buccianti@unifi.it fax: +39-055-284571

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Abstract

There are two principal chemical concepts that are important for studying the natural environment. The first one is thermodynamics, which describes whether a system is at equilibrium or can spontaneously change by chemical reactions. The second main concept is how fast chemical reactions (kinetics or rate of chemical change) take place whenever they start. In this work we examine a natural system in which both thermodynamics and kinetic factors are important in determining the abundance of NH_4^+ , NO_2^- and NO_3^- in superficial waters. Samples were collected in the Arno Basin (Tuscany, Italy), a system in which natural and antrophic effects both contribute to highly modify the chemical composition of water. Thermodynamical modelling based on the reduction-oxidation reactions involving the passage $NH_4^+ \to NO_2^- \to NO_3^-$ in equilibrium conditions has allowed to determine the E_h redox potential values able to characterise the state of each sample and, consequently, of the fluid environment from which it was drawn. Just as pH expresses the concentration of H^+ in solution, redox potential is used to express the tendency of an environment to receive or supply electrons. In this context, oxic environments, as those of river systems, are said to have a high redox potential because O_2 is available as an electron acceptor.

Principles of thermodynamics and chemical kinetics allow to obtain a model that often does not completely describe the reality of natural systems. Chemical reactions may indeed fail to achieve equilibrium because the products escape from the site of the rection or because reactions involving the trasformation are very slow, so that non-equilibrium conditions exist for long periods. Moreover, reaction rates can be sensitive to poorly understood catalytic effects or to surface effects, while variables as concentration (a large number of chemical species can coexist and interact concurrently), temperature and pressure can have large gradients in natural systems. By taking into account this, data of 91 water samples have been modelled by using statistical methodologies for compositional

data. The application of log-contrast analysis has allowed to obtain statistical parameters to be correlated with the calculated E_h values. In this way, natural conditions in which chemical equilibrium is hypothesised, as well as underlying fast reactions, are compared with those described by a stochastic approach.

Introduction

Continental freshwaters are very important to humans since they are the only reliable source of drinking water. The chemical composition of rivers, lakes and groundwaters varies widely and is governed predominantly by three factors: element chemistry, weathering regimes and biological processes (Berner and Berner, 1996). In addition, human perturbations may have a major effect (Skinner et al., 1997; Bouwman, 1998; van Breemen, 2002; Okafor and Ogbonna, 2003). In river systems Nitrogen chemistry is complex because the element is present in several oxidation states, of which N(0) nitrogen gas (N_2) , (N^{3-}) ammonium (NH_4^+) and (N^{5+}) nitrate (NO_3^-) are the most important. The nitrogen cycle is dominated by reactions involving biological material. All the reactions in the series $N_2(dinitrogen) \rightarrow NH_3(ammonia) \rightarrow NO_2^-(nitrite) \rightarrow NO_3^-(nitrate) \rightarrow aminoacids \rightarrow proteins$, and the reverse reactions back to N_2 , can be carried out by microorganisms (Krug and Winstanley, 2002).

Nitrogen gas dissolved in river water cannot be utilised as a nitrogen source by most plants and algae because they cannot break its strong triple bond. Specialised nitrogen-fixing bacteria and fungi do exist to exploit N_2 , but it is not an energetically efficient way of obtaining nitrogen. Hence, these microorganisms are only abundant when N_2 is the only available nitrogen source. Nevertheless, along with the fixation of N_2 by lightning, nitrogen-fixing microorganisms provide the major natural source of nitrogen for rivers. Biological processes use nitrogen in the 3-oxidation state, particularly as amine groups in

proteins. This is the preferred oxidation state for algal uptake and also the form in which nitrogen is released during organic matter decomposition, largely as NH_4^+ . Once released into soil or water, NH_4^+ being cationic, may be adsorbed onto negatively charged organic and inorganic coatings such as soil particles and clay mineral surfaces. Ammonium is also taken up by plants or algae, or oxidised to NO_3^- , a process that is usually catalysed by bacteria. In contrast to NH_4^+ , NO_3^- is an ionic and not retained in soils. Therefore, NO_3^- from rainwater or fertilisers or derived from the oxidation of soil organic matter and animal wastes will wash out of soils and into rivers. Apart from biological uptake, denitrification in low-oxygen environments is the most important way for which nitrate is removed from soils, rivers and groundwaters. Seasonal variations of NO_3^- concentrations in rivers from temperatue climate are caused by fluctuation in supply of NO_3^- from soil, depending on the intensity and quantity of rainfall. Increases in both the area and the intensity of agricultural activity are probably responsible for the increased NO_3^- concentrations in several rivers of Europe and North America (Codispoti and Christensen, 1985; Berner and Berner, 1996). There are entire ecosystems ranging from forests to coastal waters that are now being overhelmed by nitrogen compounds and the attention on this problem by the European Countries is increasing (Ekholm et al., 2000; Wade et al., 2002). Excess nitrate ion in wastwater flowing into seawater (for example the Baltic Sea) has resulted in algal blooms that pollute the water once they die. Excess nitrate ion in drinking water is a potential health hazard since it can result in methemoglobinemia in newborn infants as well as in adults, with a particular enzyme deficiency. Recently, an increase in the risk of acquiring non-Hodgkin's lymphoma has been found for persons consuming drinking water having high levels (long-term average of 4 ppm nitrogen as nitrate) of nitrate in drinking water. However, since epidemiological investigations to establish statistically significant positive relationship between nitrate levels in drinking water and health problems are difficult, the expenditure of public money on nitrate level reductions has become a controversial subject (Baird, 2001). Hovewer, mapping of the areas in which this type of problems is present, *i.e.* river water and ground–waters, may be important to manage the use of the land and its resources. This paper aims to give a contribution in the investigation of the abundances of NH_4^+ , NO_2^- and NO_3^- in river waters. A methodological approach based on thermodynamical and statistical bases is proposed to obtain parameters able to monitor natural systems in time and space. Data were collected in the Arno Basin (Tuscany, central–northern Italy), an area in which natural weathering phenomena and antrophical input characterise the chemical composition of the rivers. Due to the compositional nature of the collected data, their investigation has been performed in the light of the recent developments of the compositional data analysis theory (Aitchison, 1986a, b).

Geological background

The Arno River Basin, with a drainage surface of about $8,228 \ km^2$ is located within the mountain belt of the Northern Apennines (Tuscany, central-northern Italy). It originates at about $1,650 \ m$ from the SW flank of Mt. Falterona and flows for $242 \ km$ into the Tyrrhenian Sea after crossing highly urbanised, industrialised and cultivated areas. Florence, the biggest city in Tuscany, together with Arezzo, Pisa, Pistoia and Prato are the main cities contributing to major pollution. The catchment can be divided into six sub-basins from East to West: Casentino (CA), Chiana Valley (CH), Sieve (SI), Upper Valdarno (UV), Middle Valdarno (MV) and Lower Valdarno (LV) (Figure 1). The Arno Basin mainly consists of Oligocene-Miocene arenaceous-calcareous-marly flysch, structurally complex clay and Neogene marine and fluvio-lacustrine deposits (Abbate and others, 1982; Moretti, 1994). Mesozoic and Paleozoic clastic and carbonaceous rocks of the Tuscan Series (Triassic and evaporitic rocks) are exposed in the upper part of the Elsa and Era sub-basins, while Paleozoic quarzites widely occur in the Zambra sub-basin. Anthropogenic contam-

inants, such as nitrates and nitrites, in the shallow aquifers of the alluvial plain of the Middle Valdarno were associated with the infiltration of sewage effluents near the surface by Bencini and others (1993). Moreover, Bencini and others (1995) suggested that nitrate contamination largely affect the plain with nitrate concentrations as high as 159 mg/L, probably due to domestic sewage disposal of cesspools, septic tanks and leaking sewers. Conversely, a recent investigation on the Chiana Valley surface waters (Ramaldi and others, 2004) found concentrations of NH_4^+ and NO_2^- up to 18 (mean value: 2.2 mg/l) and 5.3 (mean value: 1.2 mg/l) mg/l, respectively, whereas NO_3^- concentrations are generally <10 mg/l

Starting from 2002 a wide monitoring campaign was developed with the aim to characterize the nature of the main factors influencing the chemistry of the superficial waters collected in the Arno Basin. Beside the concentrations of the major anions and cations, the abundance of NH_4^+ , NO_2^- and NO_3^- , by using spectrophotometric methodologies, was also determined for more than 450 samples. Compositional changes in the subcomposition of the nitrogen species have been modeled by both thermodynamical and statistical approach, choosing a representative group of 91 samples (Figure 1).

Methodological approach and results

The branch of physical chemistry known as thermodynamics is focused on the study of the transformations of energy and, in particular the transformation of energy from heat into work and vice versa. Thermodynamics allows to describe natural systems in terms of certain measurable properties of matter, such as volume, pressure, temperature and chemical composition of the constituent components. Geochemical modelling can be based on the fundamental laws of thermodynamics and kinetics to answer questions at the basis of everyday chemistry, such as why reactions reach equilibrium, what their composition is

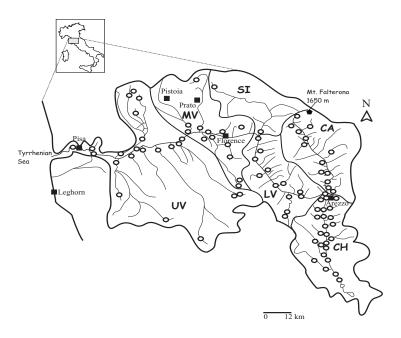


Figure 1: Map of the Arno basin with the location of the sampling sites (Tuscany, central-northern Italy): CA = Casentino, CH = Chiana Valley, SI = Sieve, UV = Upper Valdarno, MV = Middle Valdarno, LV = Lower Valdarno.

at equilibrium and so on. As mineral assemblages and fluids in nature are complicated, equilibrium thermodynamics is essential to Earth scientists and is usually the starting point for any geochemical investigation (Anderson and Crerar, 1993; Anderson, 1996). This point was established in the nineteenth century by Jacobus Van't Hoff (1852–1911) who used the Gibbs phase rule to reconstruct the composition of ancient seawater from the mineral assemblages found in evaporite deposits. Stated in terms of the primary thermodynamic functions of temperature, energy and entropy, the laws of thermodynamics can succintly be stated for any actual process as follows: first law, the total energy is conserved (i.e. remains unchanged); second law, the total entropy increases; third law, the absolute (Kelvin) temperature remains above zero degrees. When used appropriately, these laws provide the necessary and sufficient conditions for geochemists to 1) establish the equilibrium in a system that includes chemical reactions, heat transfer and mechanical

strain; 2) discharge as impossible some processes; 3) convert thermodynamic data such as the energy and entropy changes into an equilibrium constant; 4) predict the effects of changes in temperature, pressure or composition on the reaction equilibrium (Kraynov, 1997).

The most widely used geochemical modeling programs consist of a computer code plus a related file of data called *database*. The latter contains thermodynamic and kinetic parameters to be used by the code together with concentrations or other constrains as input with the aim to produce results that describe a geochemical model for a particular chemical system (Lichtner, 1996). Databases associated with geochemical modeling programs consist of a list of the basic species, and a list of secondary or auxiliary species, minerals and gases, each described in terms of the basis species, and with the equilibrium constant of the reaction linking the secondary species or mineral to the basis species. Basically the algorithms solve a set of mathematical equations describing chemical equilibria and chemical mass balance (Westall et al., 1976; Nordstrom et al., 1979; Bethke, 1996). The quality of the information reported in the data base is crucial to obtain good results and a discussion of data accuracy is reported in Nordstrom and Munoz (1986) and Criscenti et al. (1996). A summary of the most reliable codes is presented by Mangold and Tsang (1991) and Paschke and van der Heijde (1996). We have used EQ3/6 code, Version 7.2c, supported by the EQLIB library (Wolery, 1992), in order to obtain the values of E_h , the emf (electromotive force) of the oxidation-reduction reactions characterising the chemical composition of each water sample. In natural systems E_h is a parameter whose values reflect the ability of the environment to be an electron donor or acceptor relative to the standard H_2 electrode as reference. Ecosystems in contact with the atmosphere (rainwater, streams, oceans and mine waters) have positive E_h values and work as electron acceptors (oxidizing agents). On the other hand, surroundings that are isolated from the atmosphere (waterlogged soils, euxenic marine basins, organic-rich brines) have negative E_h values, and work as electron donors (reducing agents). The probability plot of the E_h values for the 91 water samples from the Arno Basin is reported in Figure 2. By comparing the continuous line related to the normal model, the probability distribution can be considered as Gaussian (Kolmogorov-Smirnov test, $\alpha = 0.05$). This means that data have been drawn from the same population and that the mean, equal to 0.6567, and the standard deviation, equal to 0.0366, are good estimators of its parameters. The E_h values reflect our samples from a thermodynamical point of view and consequently the oxidizing conditions of environments in contact with the atmosphere. The probability distribution say us that conditions more oxidant that the mean, as well as more reducent, are present with the same probability and that E_h values are given by the sum of numerous independent events.

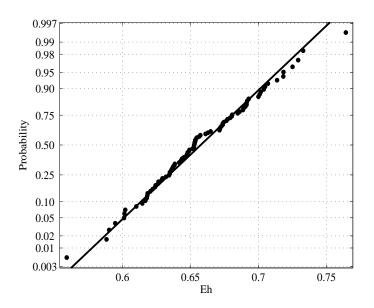


Figure 2: Normal probability plot of E_h values determined by EQ3/6 code for the Arno Basin running waters

In order to correlate this parameter with another obtained by statistical investigation, a log-contrast analysis for compositional data has been performed (Aitchison, 1983). Compositional data, consisting of vectors of proportions, are difficult to be handled statisti-

cally because of the awkward constraint that the components of each vector must sum to unity. Such data sets frequently display marked curvature so that linear techniques such as standard principal component analysis are likely to prove inadequate. Aitchison (1981; 1982) has proposed a resolution of the constrain difficulty in the analysis of compositinal data. In practical terms this consists of first trasforming each compositional vector $x^{(d)} = (x_1, \dots, x_d)$ in S^{d-1} (the simplex space) to a vector $y^{(d-1)} = (y_1, \dots, y_{d-1})$ in R^{d-1} . The log ratio transformation $y^{(d-1)} = \ln(x_{-j}/x_j)$ is then applied, where x_{-j} denotes the vector $x^{(d)}$ with x_j omitted. Transformed vectors are finally analysed by standard multivariate methods available in R^{d-1} . The nonlinearity of the logarithmic function opens up the possibility to capture the curvature in data sets and its approximate linerity over parts of its range also makes feasible the modelling of linear data sets. However, in this case a major difficult remains since different choices of the divisor x_j lead to a different principal components. We can solve this problem by working with a more symmetric approach if the denominator is given by the geometric mean of the d components of the composition $g(x) = (x_1, \dots, x_d)^{1/d}$. The covariance matrix that is investigated in this case is given by:

$$\Omega = cov\{\ln(x^{(d)}/g(x))\}. \tag{1}$$

The matrix Ω is positive-semidefinite, its one zero eigenvalue having an associated eigenvector u_d consisting of d units. The other d-1 eigenvalues $\lambda_1, \dots, \lambda_{d-1}$, labelled in descending order of magnitude, are positive and the corresponding eigenvectors a_1, \dots, a_{d-1} , being orthogonal to u_d , yield log linear contrasts of the proportions. They are expressed as:

$$\sum_{i=1}^{d} a_i \ln x_i = \sum_{i=1}^{d} a_i \ln(x_i/g(x)), \tag{2}$$

where $a_1, \dots, a_d = 0$. The first log constrast, explaining about the 74.8% of the data variability, is given by

$$-0.39\ln(NH_4^+) - 0.43\ln(NO_2^-) + 0.81\ln(NO_3^-) = \kappa_1. \tag{3}$$

If the equation is semplified and written in exponential form we obtain

$$\frac{(NO_3^-)}{(NH_4^+)^{0.5} \cdot (NO_2^-)^{0.5}} = exp(\kappa_1),\tag{4}$$

leading to the relationship NH_4^+ : NO_2^- : $NO_3^- = 1$: 1: 2. The κ_1 variable follows the Gaussian curve (Kolmogorov-Smirnov test, $\alpha = 0.05$). Its normal probability plot is reported in Figure 3. The second log constrast, explaining about the 25.2% of the data variability is given by:

$$0.72\ln(NH_4^+) - 0.69\ln(NO_2^-) - 0.02\ln(NO_3^-) = \kappa_2, \tag{5}$$

or semplifying:

$$\frac{(NH_4^+)}{(NO_2^-)} \sim exp(\kappa_2),\tag{6}$$

due to the low value of NO_3^- coefficient. In this case the relationship $NH_4^+:NO_2^-=1:1$ is obtained. The κ_2 variable is again descrivible with the Gaussian model (Kolmogorv-Smirnov test, $\alpha=0.05$) and its probability plot is reported in Figure 4.

The two log constrasts describing relationships among the variables of the sub composition $NH_4^+ - NO_2^- - NO_3^-$ can be reported on a ternary diagram as in Figure 5. Another way to see the reciprocal weight of the components of the composition is the variation diagram as those reported in Figure 6. The compositional lines represent the reciprocal relationships among the members of the composition as described by the first log contrast. The linear

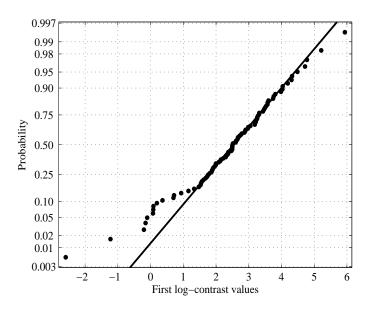


Figure 3: Normal probability plot of the first log-contrast values

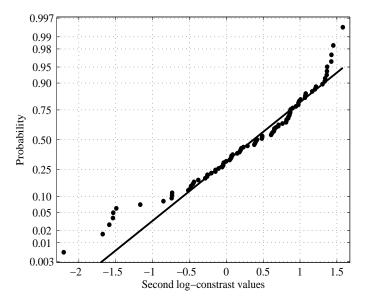


Figure 4: Normal probability plot of the second log-contrast values

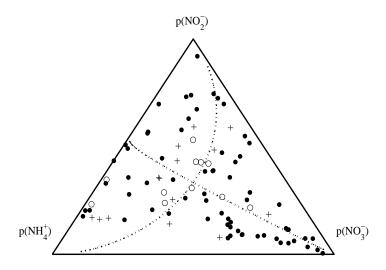


Figure 5: Ternary diagrams of the system NH_4^+ - NO_2^- - NO_3^- (+ = Chiana, \circ = Canale Maestro, \bullet = Arno valley); data have been perturbed by the vector $\mathbf{p} = [35.9161.672.41]$ to obtain a better visualisation (the original barycentre of the data moved toward the barycentre of the triangle).

patterns of Figures 5 and 6 are obtained by using perturbation and power transformation operations, both able to give to the simplex a (d-1) dimensional vector space structure (Barceló-Vidal et al., 2001; Pawlowsky-Glahn and Egozcue, 2002). They are given by $\alpha \odot \mathbf{p}$ where \mathbf{p} is the unitary perturbation vector obtained from log contrast analysis, α is a scalar value that ranges from -3 to +3 and \odot indicates the power transformation operation. The α values needed to post observations on the plot can be obtained from the log contrast scores or by projecting each observation onto the linear trend given by the log contrast (Pawlowsky-Glahn and Egozcue, 2001). The behaviour of NO_3^- is well represented by the model for all the α values while those of NH_4^+ and NO_2^- show a wide scattering of the real data around the models, particularly far from the barycentre (0 value), towards negative values. The κ_1 values have been correlated with the E_h values as obtained by the thermodynamical modelling. The results of the linear fitting are reported

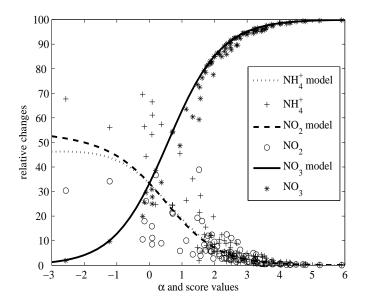


Figure 6: Relative variation diagram of the system NH_4^+ - NO_2^- - NO_3^- as modelled by $\alpha \odot \mathbf{p}$; α is a scalar value that ranges from -3 to +3 while \mathbf{p} is the unitary perturbation vector obtained from log contrast analysis.

in Figure 7. In Figure 8 the bar plot of the residulas is also reported. Notwithstanding the scarse goodness of the linear fitting (correlation between the variables is 0.44), the line can be considered as a reference to characterise data more and more far from it. The bar plot of Figure 8 help us to visualise the difference between the E_h calculated values and those obtained by the model. If only the n=59 samples inside the area between $-0.03 \div +0.03$ (limits corresponding to the first and third quartile) are considered, the correlation among the two variables move from 0.44 to 0.70. These samples are characterised by electrical conductivity ranging in the interval $0.12 \div 6.6$ mS/cm (median 0.84 mS/cm) and nitrogen species in the intervals $0.026 \div 6.32$ mg/l for NH_4^+ (median 0.22), $0.01 \div 7.28$ mg/l for NO_2^- (median 0.20), $0.10 \div 32.5$ mg/l for NO_3^- (median 32.5). All the samples out from the area are characterised for the lower values of conductivity $(0.22 \div 1.86$ mS/cm, median 0.73 ms/cm) as well as of NH_4^+ ($0.040 \div 4.39$ mg/l, median 9×10^{-2}), NO_2^- ($0.01 \div 1.97$

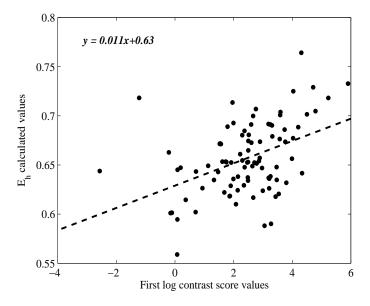


Figure 7: Linear relationship between the first log contrast values and the E_h calculated data.

mg/l, median 6.39×10^{-2}) and NO_3^- (0.12÷30.5 mg/l, median 2.33). In Figure 9 data have been mapped by considering the discrimination in two groups, according to the results of Figure 8. Full black circles are related to samples pertaining to difference values between the calculated and the modelled E_h in the range $-0.03 \div +0.03$. For them an equilibrium condition is hypothesised and deduced by the first log contrast k_1 values. On the other hand, open circles are associated to samples characterised by high negative or positive values of the difference. On the whole these samples are related to situations characterised by the lowest values of nitrogen species and conductivity of the investigated area.

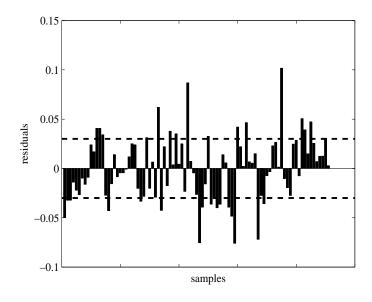


Figure 8: Bar plot of the residuals of the linear fitting.

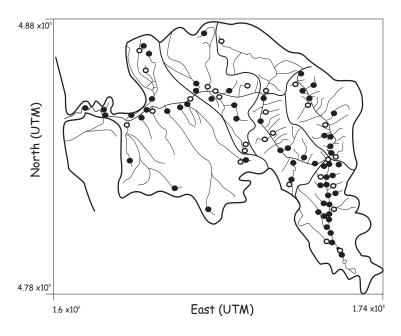


Figure 9: Spatial variation of the difference values between the calculated and the modelled E_h . Full black circles are related to samples pertaining to difference values in the range $-0.03 \div +0.03$; open circles are associated to samples characterised by high negative or positive values of the difference.

Discussion

The redox potential of a natural system is determined by the physico-chemical features of the system itself. Geochemists use $E_h - pH$ diagrams to predict the oxidation state of various constituents in natural environments. E_h determines what modes of microbial activity are possible in a given environment. Few oxic environments have redox potential of less than +0.6 V and a progressive decrease in redox potential occurs when soils are flooded. Redox potential drops as heterotrophic respiration of organic carbon depletes the soil of O_2 and in most cases, organic matter contributes a large amount of reducing power that lowers the redox potential (Bartlett, 1986). The results of many studies suggest that a particular sequence of reactions is expected, as progressively lower redox potentials are achieved (Achtnich et al., 1995; Lovley, 1995; Peters and Conrad, 1996). In our case the E_h calculated values (EQ3/6 code) ranges in the interval 0.559÷0.764 V with the mean representive value of 0.6567 V, and from a thermodynamical point of view describe chemical equilibrium conditions. The fact that E_h values are modelled by a probability density function as the Gaussian reveals a sort of statistical equilibrium. Values higher or lower with respect to the mean are expected with the same probability and represent the fluctuations due to the sum of several independent factors affecting the oxidation-reduction conditions of the investigate ecosystem. In order to link thermodynamical and statistical modelling, E_h values have been correlated with the κ_1 scores of the first log-contrast (74.8% of the data variability). On the whole they represent the relationship $NH_4^+:NO_2^-:NO_3^-=1:1:2$, quantitatively shifted towards NO_3^- abundance, the final species of the oxidation processes. The linear relationship between E_h and κ_1 is not so good due to the wide scattering of the data. Apparently this result may be attributable to the behavior of κ_1 (Figure 3) showing some anomalous values in the tails of the frequency distribution. Thus, it is evident that the variability described by log constrast analysis it is not equivalently revealed by thermodynamical modelling. The distance value of each sample from the best fitted linear model in the plot $\kappa_1 - E_h$ has been used to point out cases where correlation is susteinable. Cases for which the distance values are in the range $-0.03 \div +0.03$ show a correlation between κ_1 and E_h equal to 0.70. In this case, the relationships among the variables are not quantitatively disturbed for absence (scarse abundance) of one of the components. All the observations pertain to the right part of the variation diagram of Figure 6, thus indicating that the system is fastly moving towards NO_3^- .

Cases for which the linear relationship between κ_1 and E_h is not a validate model represent situations of discordance between thermodynamical and statistical results. In this case observations are characterised by low values of NH_4^+ , NO_2^- , NO_3^- , as well as of conductivity, all indications of low recent and ancient pollution levels. Data pertaining to the two groups (κ_1 and E_h correlated or uncorrelated) have been mapped and the results indicate us that different conditions (chemical equilibrium described or not described by log contrast analysis) can closely be associated in space. The system $NH_4^+ - NO_2^- - NO_3^-$ is thus dominated by the presence of punctual phenomena with low spatial correlation overlapped to a general situation of chemical equilibrium. This circumstance can lead, in general, to serious problems in the mapping of interpolated values. A way to solve this question may be the plotting of data, for which thermodynamical equilibrium is also correspondent to the statistical description of the system, separately. Equilibrium data can reasonably represent a continuos phenomenon in space and interpolation may have sense.

Conclusions

Most urban and agricultural-zootechnical areas are affected by the problem of nitrate pollution in both surface and ground waters. Despite the increasing efforts to reduce NO_3^-

inputs from intensive agriculture at national and European (EC Directive 91/676/CEE) levels, nitrate is still one of the main contaminants. Many aquatic environments have elevated NO_3^- levels as the result of anthropogenic activities that involve nitrogen compounds deriving from mineral fertilizers, septic systems, animal manures and so forth. Development of both effective management practices to preserve water quality and remediation plans for contaminated sites require the identification of pollutant sources and consequently of the processes affecting local NO_3^- concentrations. In this framework the combination of thermodynamical and statistical modelling may represent a powerful tool to describe natural ecosystems and to map phenomena showing a spatial continuity with respect to those dominated by punctual factors. E_h values determined for waters collected in the Arno basin by EQ3/6 code, under the reasonable hypothesis of chemical equilibrium, have been matched with the scores of the first log constrast summarising the reciprocal relationships among NH_4^+ , NO_2^- and NO_3^- . Results indicate that the approach can be used to discriminate data for which the concept of chemical equilibrium is associated with the statistical modelling of the variability, thus identifying potentially continuos spatial phenomena. These samples are closely associated in space with those characterised by local high variability and low spatial correlation. If sets of data with these features are considered together, the interpretation of maps obtained by interpolation procedures can be, in general, highly compromised.

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